

# Effects of moisture and carbonate additions on CO<sub>2</sub> emission from calcareous soil during closed-jar incubation

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**Abstract:** Calcareous soil contains organic and inorganic carbon (C) pools, which both contribute to CO<sub>2</sub> emission during closed-jar incubation. The mineralization of organic C and dissolution of inorganic C are both related to soil moisture, but the exact effect of water content on CO<sub>2</sub> emission from calcareous soil is unclear. The objective of this experiment was to determine the effect of soil water content (air-dried, 30%, 70%, and 100% water-holding capacity (WHC)), carbonate type (CaCO<sub>3</sub> or MgCO<sub>3</sub>), and carbonate amount (0.0, 1.0%, and 2.0%) on CO<sub>2</sub> emission from calcareous soil during closed-jar incubation. Soil CO<sub>2</sub> emission increased significantly as the water content increased to 70% WHC, regardless of whether or not the soil was amended with carbonates. Soil CO<sub>2</sub> emission remained the same or increased slowly as the soil water content increased from 70% WHC to 100% WHC. When the water content was ≤30% WHC, soil CO<sub>2</sub> emission from soil amended with 1.0% inorganic C was greater than that from unamended soil. When the soil water content was 70% or 100% WHC, CO<sub>2</sub> emission from CaCO<sub>3</sub> amended soil was greater than that from the control. Furthermore, CO<sub>2</sub> emission from soil amended with 2.0% CaCO<sub>3</sub> was greater than that from soil amended with 1.0% CaCO<sub>3</sub>. Soil CO<sub>2</sub> emission was higher in the MgCO<sub>3</sub> amended soil than from the unamended soil. Soil CO<sub>2</sub> emission decreased as the MgCO<sub>3</sub> content increased. Cumulative CO<sub>2</sub> emission was 3–6 times higher from MgCO<sub>3</sub> amended soil than from CaCO<sub>3</sub> amended soil. There was significant interaction effect between soil moisture and carbonates on CO<sub>2</sub> emission. Soil moisture plays an important role in CO<sub>2</sub> emission from calcareous soil because it affects both biotic and abiotic processes during the closed-jar incubation.

**Keywords:** calcareous soil; soil moisture; organic carbon; CO<sub>2</sub> emission

**Citation:** YanJie DONG, Miao CAI, JianBin ZHOU. 2014. Effects of moisture and carbonate additions on CO<sub>2</sub> emission from calcareous soil during closed-jar incubation. *Journal of Arid Land*, 6(1): 37–43. doi: 10.1007/s40333-013-0195-6

Soil is both a large carbon (C) sink and a large C source in the terrestrial ecosystem (Schlesinger and Andrews, 2000). Soil respiration releases 68–80 Pg C/a, the second largest C flux between the terrestrial ecosystem and the atmosphere (Raich et al., 2002). Even a small change in soil respiration could significantly intensify or mitigate current atmospheric increases of CO<sub>2</sub>, with potential feedbacks to climate change (Freer-Smith et al., 2008).

Most CO<sub>2</sub> released by soil respiration is from the mineralization of soil organic C. Mineralization is primarily a biotic process, but it is also influenced by environmental factors, especially soil moisture (Leirós et al., 1999). Closed-jar incubation is a common method for estimating soil respiration (Alef, 1995). The closed-jar method is widely used to quantify soil microbial activity and to investigate the effects of precipitation change on CO<sub>2</sub> emission from different soil types

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Received 2012-12-27; revised 2013-03-11; accepted 2013-04-07

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(Borken et al., 2003; Rey et al., 2005; Inglema et al., 2009; Merbold et al., 2009; Muhr et al., 2010). Most studies have focused on CO<sub>2</sub> emission from acidic or neutral soils. Few researchers have studied the effect of soil moisture on CO<sub>2</sub> emission from calcareous soil.

Calcareous soil, which covers more than 30% of the Earth's land surface (Chen and Barak, 1982), contains two pools of C, i.e. organic C and inorganic C. The inorganic C content of calcareous soil is typically 2–5 times greater than the organic C content. In some arid regions, inorganic C content can be ten times greater than the organic C content (Schlesinger, 1982). Some researchers have suggested that carbonates could dissolve and release CO<sub>2</sub> during incubation, leading to an overestimation of organic C mineralization in soil (Alef, 1995; Bertrand et al., 2007). The dissolution of CaCO<sub>3</sub> is a chemical process that is closely related to soil water content (Birkeland, 1984). Thus, the biotic and abiotic processes that contribute to soil CO<sub>2</sub> emission are both influenced by soil moisture. The exact effect of water content on CO<sub>2</sub> emission from calcareous soil is unclear.

There are two types of carbonates in calcareous soil: CaCO<sub>3</sub> and MgCO<sub>3</sub> (Guo, 1992). Due to its higher dissolubility, MgCO<sub>3</sub> is easily dissolved in the soil and leached. The content of MgCO<sub>3</sub> is less than 10% in calcareous soil. In contrast, low dissolubility contributes to the easy accumulation and conservation of CaCO<sub>3</sub> in the soil, making it the main composition of soil carbonate. Despite its low content in calcareous soil, MgCO<sub>3</sub> is more active than CaCO<sub>3</sub>. So when we discuss the stabilization of soil carbonates, the influence of MgCO<sub>3</sub> should not be ignored.

Global warming is predicted to increase the number of abnormal rainfall events (IPCC, 2001). Therefore, it is important to understand the effect of soil moisture on CO<sub>2</sub> emission from calcareous soil. The aims of this study are (1) to evaluate the effect of soil moisture on CO<sub>2</sub> emission from calcareous soil during closed-jar incubation, and (2) to estimate the effect of soil carbonate type and amount on CO<sub>2</sub> emission during incubation.

## 1 Materials and methods

### 1.1 Study area

Soil samples were collected from the Experiment Sta-

tion at Northwest A&F University, Yangling, Shaanxi province, China (34°17'56"N, 108°04'07"E). The elevation at the site is 523 m, the annual average temperature is 13°C, and the average annual precipitation is 600 to 650 mm. The soil type is Lou soil, which is classified as an Eum-Orthric Anthrosol (equivalent to a Udic Haplustalf in the USDA system). Lou soil has been cultivated for more than two thousand years. The inorganic C content in the 0–200 cm soil profile ranged between 0.7 and 22.8 g/kg.

### 1.2 Soil samples

The soil samples in this study were from the eluvial horizon (60–80 cm depth) of the profile. We chose this depth in order to minimize the effects of preexisting carbonates on CO<sub>2</sub> emission during incubation. The inorganic C content of the eluvial horizon was 0.74 g/kg. The soil was air-dried, passed through a 2-mm sieve, and then stored.

The physical and chemical properties of the soil were measured using standard methods. Soil organic C was measured with the Walkley-Black acid digestion method, inorganic C with the aerometric method, total N with the Kjeldahl method, and soil pH with a pH meter (soil:water=1:5). The soil organic C content was 5.81 g/kg, the total nitrogen (N) content was 0.24 g/kg, and the soil pH was 7.67.

### 1.3 Incubation method

To investigate the effects of soil moisture and carbonate addition on CO<sub>2</sub> emission during closed-jar incubation, we designed an experiment that included four soil moisture contents (air-dried, 30%, 70% and 100% of field water holding capacity (WHC)) and three carbonate concentrations (0, 1.0% and 2.0% inorganic C added as either CaCO<sub>3</sub> or MgCO<sub>3</sub>). Twenty grams of air-dried soil were placed in 50-mL plastic bottles. Deionized water was added to adjust the water content and then CaCO<sub>3</sub> or MgCO<sub>3</sub> was added to the soil. The samples were put into 500-mL jars along with bottles containing 10-mL 0.1 mol/L NaOH. The jars were sealed and then incubated at 25°C. Soil moisture was kept constant throughout the incubation by weighing the bottles every three days and adding deionized water when necessary. The NaOH solutions were changed on day 1, 2, 4, 8, 12, 16, 23, 30, 40, 51, 61,

67 and 83. The incubation time was chosen because cumulative CO<sub>2</sub> emission would remain unchanged after a 83-d incubation. And the jars were sealed during the incubation except adding water every three days. The amount of CO<sub>2</sub> trapped in the NaOH was measured with the HCl-BaCl<sub>2</sub> titration method (Alvarez et al., 1995). Each treatment was replicated three times.

#### 1.4 Statistic analysis

CO<sub>2</sub> emission during the interval  $i$  ( $C_i$ ) was calculated as Eq. 1 (Chen et al., 2009):

$$C_i = \frac{(V_s - V_a) \times n \times 12 \times 1000}{20 \times 2}. \quad (1)$$

Where  $V$  is the volume of HCl used to neutralize the NaOH solution,  $s$  stands for the soil samples,  $a$  stands for the air in the jar, and  $n$  stands for the molar concentration of HCl solution.

The cumulative CO<sub>2</sub> emission ( $C$ ) during the incubation was calculated as Eq. 2:

$$C = \sum C_i. \quad (2)$$

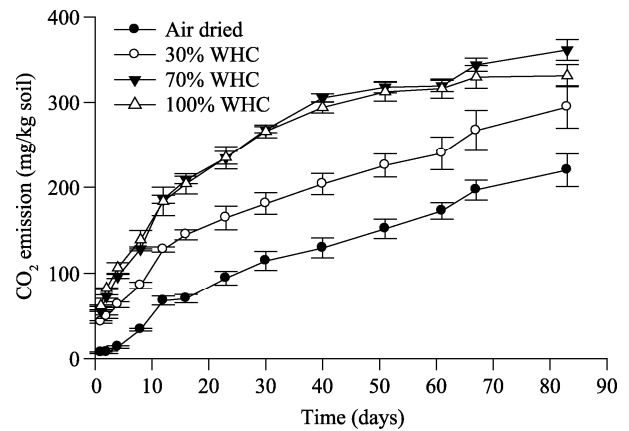
Cumulative CO<sub>2</sub> emission was calculated for each treatment combination using mean CO<sub>2</sub> mineralization rates ( $n=3$ ). Soil CO<sub>2</sub> emission rates were analyzed by analysis of variance (ANOVA) using the GLM procedure of SAS (SAS Institute Inc., 2000). When the effects of soil moisture or C content were significant ( $P<0.05$ ), paired  $t$ -tests were used to compare CO<sub>2</sub>-C emission between the treatments. The interaction effect between soil moisture and inorganic C content on CO<sub>2</sub> emission was analyzed by ANOVA. The relationship between soil moisture and CO<sub>2</sub> emission was determined using nonlinear regression analysis (Sigmaplot 12.0, 2012). The values used in the regression analyses were the mean values of each treatment.

## 2 Results

### 2.1 Effect of moisture on soil CO<sub>2</sub> emission

Cumulative CO<sub>2</sub> emission increased significantly as the soil water content increased from air-dried to 70% WHC ( $P<0.01$ ). There was generally no significant difference in cumulative CO<sub>2</sub> emission between the 70% WHC treatment and the 100% WHC treatment (Fig. 1). The emission of CO<sub>2</sub> was 63% and 30% higher in the 30% and 70% WHC treatments respec-

tively than in the air-dried treatment ( $P<0.01$ ). Cumulative CO<sub>2</sub> emission was 6.51% less in the 70% WHC treatment than in the 100% WHC treatment.

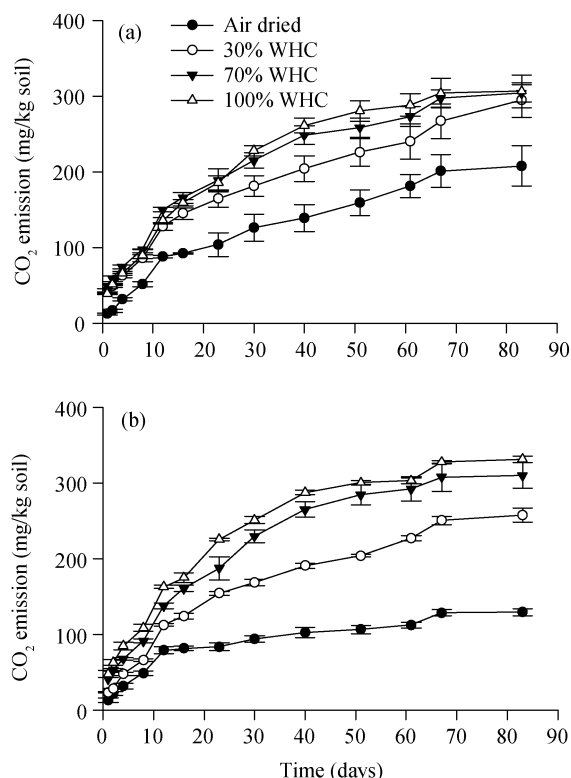


**Fig. 1** Cumulative CO<sub>2</sub> emission from soil without carbonate amendment. Error bars represent the standard error of the mean. WHC, water holding capacity.  $n=3$ .

### 2.2 Effects of moisture and CaCO<sub>3</sub> addition on soil CO<sub>2</sub> emission

Cumulative CO<sub>2</sub> emission increased significantly as soil moisture increased in the CaCO<sub>3</sub> amended soil ( $P<0.01$ ) (Fig. 2). In the 1.0% CaCO<sub>3</sub> treatment, CO<sub>2</sub> emissions from the wetted soils (i.e. 30%, 70%, and 100% WHC) were 53% to 78% higher than those from the air-dried soil. In the 2.0% CaCO<sub>3</sub> treatment, CO<sub>2</sub> emissions from the wetted soils were 90% to 334% higher than those from the air-dried soil. There was a significant quadratic relationship between cumulative CO<sub>2</sub> emission ( $y$ ) and soil moisture ( $x$ ) ( $y = -0.0208x^2 + 4.0689x + 0.0071$ ,  $R^2=0.999$ ).

The effect of CaCO<sub>3</sub> addition on CO<sub>2</sub> emission varied depending on soil water content (Fig. 2). The interaction effect between soil moisture and CaCO<sub>3</sub> concentration on CO<sub>2</sub> emission was significant ( $P<0.01$ ). When soil moisture was  $\leq 70\%$  WHC, CO<sub>2</sub> emission was less in the CaCO<sub>3</sub> amended soil than in the unamended soil. Furthermore, CO<sub>2</sub> production decreased as the amount of CaCO<sub>3</sub> added to the soil increased. However, when soil moisture was 100% WHC, CO<sub>2</sub> emission was greater in the CaCO<sub>3</sub> amended soil than in the unamended soil. Furthermore, CO<sub>2</sub> production increased as the CaCO<sub>3</sub> content increased.

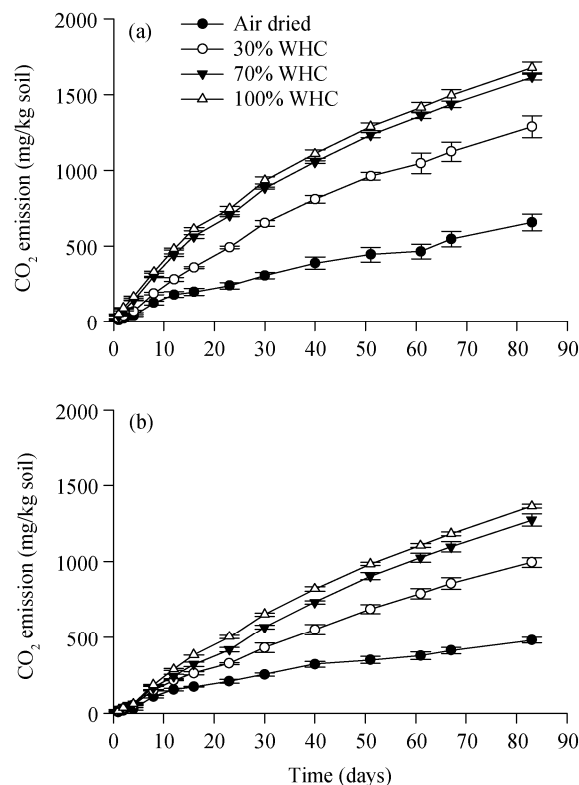


**Fig. 2** CO<sub>2</sub> emission from soil amended with 1.0% CaCO<sub>3</sub> (a) and 2.0% CaCO<sub>3</sub> (b). Error bars represent the standard error of the mean. WHC, water holding capacity.  $n=3$ .

### 2.3 Effects of moisture and MgCO<sub>3</sub> addition on soil CO<sub>2</sub> emission

Cumulative CO<sub>2</sub> emission increased significantly as soil moisture increased in the MgCO<sub>3</sub> amended soil ( $P<0.05$ ) (Fig. 3). In the 1.0% MgCO<sub>3</sub> treatment, CO<sub>2</sub> emissions from the wetted soils (i.e. 30%, 70%, and 100% WHC) were 96% to 155% higher than those from the air-dried soil. A quadratic relationship existed between cumulative CO<sub>2</sub> emission and soil moisture ( $y=-0.1011x^2+18.675x+12.959$ ,  $R^2=0.995$ ). In the 2.0% MgCO<sub>3</sub> treatment, CO<sub>2</sub> emissions from the wetted soils were 106% to 183% higher than those from the air-dried soil. In soil amended with 2.0% MgCO<sub>3</sub>, CO<sub>2</sub> emissions were 183% higher in the 100% WHC treatment, 164% higher in the 70% WHC treatment, and 106% higher in the 30% WHC treatment than in the air-dried soil. There was also a significant quadratic relationship between the cumulative CO<sub>2</sub> emission and soil moisture ( $y=-0.1353x^2+23.48x+13.865$ ,  $R^2=0.996$ ).

Cumulative emissions were highest in the 2.0% MgCO<sub>3</sub> treatment and lowest in the non MgCO<sub>3</sub> treatment (Fig. 3). The trend was different for the CaCO<sub>3</sub> amended soil. There was an exponential relationship between cumulative CO<sub>2</sub> emission and the amount of MgCO<sub>3</sub> added to the soil at different soil moisture contents. The interaction effect between soil moisture and MgCO<sub>3</sub> on CO<sub>2</sub> emission was also significant ( $P<0.01$ ).



**Fig. 3** CO<sub>2</sub> emission from soil amended with 1.0% MgCO<sub>3</sub> (a) and 2.0% MgCO<sub>3</sub> (b). Error bars represent the standard error of the mean. WHC, water holding capacity.  $n=3$ .

## 3 Discussion

### 3.1 Soil moisture and CO<sub>2</sub> emission

CO<sub>2</sub> emitted from soil during this incubation experiment originated either from organic matter or from inorganic carbon. The mineralization of organic matter is mainly controlled by the biological activities in soil and is affected by environmental factors, such as soil temperature and moisture (Rey et al., 2005; Jia et al., 2006; Wang et al., 2010). The release of CO<sub>2</sub> from inorganic sources is a chemical process and is mostly

affected by environmental factors (Ryskov et al., 2008). This study showed a significant relationship between CO<sub>2</sub> emission and soil moisture, regardless of whether or not the soil was amended with carbonates. This indicates that soil water content affects both the biotic and abiotic processes that result in CO<sub>2</sub> emission.

Soil CO<sub>2</sub> emission increased when the soil water content was  $\leq 70\%$  WHC, regardless of whether or not the soil was amended with carbonates. When the soil moisture content was  $>70\%$  WHC, the CO<sub>2</sub> emission remained the same or increased slowly as the water content increased. Other studies showed that CO<sub>2</sub> emission increased as the water content increased in peat soil (Wang et al., 2010), paddy soil (Zhang et al., 2007) and forest soil (Bowden et al., 1998). However, CO<sub>2</sub> emission decreased when soil water content exceeded 60% or 70% WHC. The significant interaction effect between soil water content and CO<sub>2</sub> emission reflects the fact that soil water content affects both the biotic and abiotic processes which result in CO<sub>2</sub> emission. When soil moisture is low, microbial activities are inhibited and the dissolution of inorganic C is reduced. Both of these factors contribute to a reduction in CO<sub>2</sub> emission. Microbial activities increase as soil moisture increases. Soil CO<sub>2</sub> emission reached a peak when soil water content was between 50% and 80% WHC (Raich and Potter, 1995). The dissolution of inorganic C also accelerated as the soil water content increased, leading to an additional increase in soil CO<sub>2</sub> emission. The concentration of O<sub>2</sub> in the soil declines as the soil water content increases to above 80% WHC. Near anaerobic conditions could have inhibited microbial activities and reduced CO<sub>2</sub> emission (Raich and Potter, 1995). Therefore, when evaluating the effect of soil moisture on CO<sub>2</sub> emission from calcareous soil, both biotic and abiotic processes should be considered.

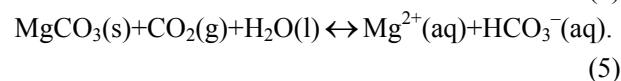
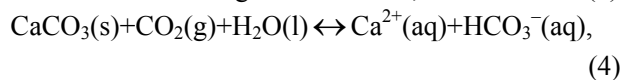
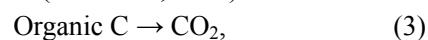
We also observed that there was a decreasing trend of CO<sub>2</sub> emission from unamended soil when soil moisture increased from 70% to 100% WHC, whereas CO<sub>2</sub> emission from soil amended with either CaCO<sub>3</sub> or MgCO<sub>3</sub> increased. The most likely reason for this is that although biological activities were reduced at 100% WHC, the dissolution of inorganic C was pro-

moted when the content of carbonates in soil was high.

### 3.2 Forms of carbonates and CO<sub>2</sub> emission

This study indicated that the effect of additional inorganic C on total CO<sub>2</sub> emission during closed-jar incubation varied with soil moisture. When the soil water content was  $<30\%$ , soil CO<sub>2</sub> emissions were higher in the 1.0% CaCO<sub>3</sub> treatment than in the control treatment. In contrast, CO<sub>2</sub> emissions were lower in the 2.0% CaCO<sub>3</sub> amendment than in the control treatment. However, when the water content reached 70% or 100% WHC, CO<sub>2</sub> emission in the CaCO<sub>3</sub> amended soil was higher than in the control treatment. Soil CO<sub>2</sub> emission in the soil increased with the increase of CaCO<sub>3</sub> content.

Soil CO<sub>2</sub> emission was greater in the MgCO<sub>3</sub> amended soil than in the control treatment. Furthermore, CO<sub>2</sub> emission decreased as the carbonate content of the soil increased. Soil CO<sub>2</sub> emissions were 3 to 6 times greater in MgCO<sub>3</sub> amended soil than in the CaCO<sub>3</sub> amended soil. These results could be explained by the dissolution and formation of carbonates in the following equations (Birkeland, 1984):



The partial pressure of CO<sub>2</sub> and the soil water content are two factors affecting the dissolution of carbonates. When the soil water content is  $<30\%$  WHC, soil microbial activities are low and the amount of CO<sub>2</sub> derived from soil organic C is relatively small. Under this condition, carbonates in the 1.0% CaCO<sub>3</sub> treatment would dissolve quickly and reach an equilibrium. However, it would take longer for the system to equilibrate in the 2.0% CaCO<sub>3</sub> treatment. Microbial activities are higher at soil moisture contents of 70% and 100% WHC than at 30% WHC. Under these conditions, microbial activities would produce enough CO<sub>2</sub> to complete the balance of Eq. 4. The CaCO<sub>3</sub> could dissolve and release CO<sub>2</sub> at these two levels of soil moisture. Soil CO<sub>2</sub> emission increased as the amount of CaCO<sub>3</sub> added to the soil increased whereas CO<sub>2</sub> emission decreased as the amount of MgCO<sub>3</sub>

added to the soil increased. One explanation is that  $\text{MgCO}_3$  dissolves more readily than  $\text{CaCO}_3$  (Patnaik, 2003). At low soil moisture,  $\text{MgCO}_3$  can complete the balance of Eq 5.

In this experiment, we studied the effect of soil water content on  $\text{CO}_2$  emission from calcareous soil during closed-jar incubation. It must be emphasized that the incubation conditions in the laboratory are different from those in the field. The dissolution and movement of carbonates in soil are more complex under field conditions. During the rainy season, carbonates in soil are leached, accumulating in the lower depths of the soil profile. Furthermore,  $\text{Ca}^{2+}$  in the soil can react with  $\text{HCO}_3^-$  produced by plants to form heavy calcium carbonates (Guo, 1992). Therefore, seasonal variation of precipitation should be taken into consideration when evaluating the effect of soil water content on  $\text{CO}_2$  emission in the field.

## 4 Conclusions

In summary, soil moisture had a significant effect on  $\text{CO}_2$  emission from calcareous soil during closed-jar incubation. Furthermore soil moisture and carbonate content had significant interaction effects on  $\text{CO}_2$  emission. Therefore, when determining  $\text{CO}_2$  emission from calcareous soil, the contribution of  $\text{CO}_2$  from carbonates needs to be considered, especially when soil moisture is not constant. In this study, we could not distinguish between  $\text{CO}_2$  released by biotic processes and that released by abiotic processes. The  $\delta^{13}\text{C}$  method can be used to solve this problem in future studies.

## Acknowledgments

This work was supported by the National Natural Science Foundation of China (40773057) and the National Technology R&D Pillar Program in the 12<sup>th</sup> Five-Year Plan of China (2012BAD15B04). We are grateful to Dr. Jeff GALE for language editing.

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